

Relating hygroscopicity and composition of organic aerosol particulate matter

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Abstract. A hygroscopicity tandem differential mobility analyzer (HTDMA) was used to measure the water uptake (hygroscopicity) of secondary organic aerosol (SOA) formed during the chemical and photochemical oxidation of several organic precursors in a smog chamber. Electron ionization mass spectra of the non-refractory submicron aerosol were simultaneously determined with an aerosol mass spectrometer (AMS), and correlations between the two different signals were investigated. SOA hygroscopicity was found to strongly correlate with the relative abundance of the ion signal m/z 44 expressed as a fraction of total organic signal (f_{44}). m/z 44 is due mostly to the ion fragment CO_2^+ for all types of SOA systems studied, and has been previously shown to strongly correlate with organic O/C for ambient and chamber OA. The analysis was also performed on ambient OA from two field experiments at the remote site Jungfraujoch, and the megacity Mexico City, where similar results were found. A simple empirical linear relation between the hygroscopicity of OA at subsaturated RH, as given by the hygroscopic growth factor (GF) or “ κ_{org} ” parameter, and f_{44} was determined and is given by $\kappa_{\text{org}} = 2.2 \times f_{44} - 0.13$. This approximation can be further verified and refined as the

database for AMS and HTDMA measurements is constantly being expanded around the world. The use of this approximation could introduce an important simplification in the parameterization of hygroscopicity of OA in atmospheric models, since f_{44} is correlated with the photochemical age of an air mass.

1 Introduction

The hygroscopic growth of atmospheric aerosol particles plays an important role in numerous atmospheric processes such as climate forcing, visibility degradation, cloud formation and heterogeneous chemistry. Atmospheric aerosol components can be classified into inorganic and organic fractions. Organic material, habitually referred to as organic aerosol (OA), often represents more than half of the mass for submicron particles (Jimenez et al., 2009). A substantial fraction of that is secondary OA (SOA) which is formed from chemical reactions of gaseous compounds (Zhang et al., 2007a). The hygroscopic properties of most inorganic salts present in atmospheric aerosol are well known. In contrast, OA is composed of thousands of individual species (Goldstein and Galbally, 2007). Of these, the hygroscopic properties have been investigated only for very few substances. Different approaches are used to determine the influence of OA



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on the hygroscopic behavior of ambient aerosol under subsaturated relative humidity (RH) conditions. Previous laboratory studies, investigating the hygroscopic properties of organic model substances, have mostly focused on pure and mixed low molecular weight organic acids such as carboxylic acids, dicarboxylic acids, and multifunctional organic acids or their salts (Peng et al., 2001; Choi and Chan, 2002b), as well as on their mixtures with inorganic salts (Cruz and Pandis, 2000; Choi and Chan, 2002a; Hämeri et al., 2002). Theoretical models for the prediction of the hygroscopic growth of mixed organic and mixed inorganic/organic aerosols were recently introduced by several investigators (Clegg et al., 2001; Topping et al., 2005b). Treating individual compounds at the molecular level requires binary aqueous data for each compound (Saxena and Hildemann, 1997). The problem of characterizing activity coefficients in multiple-component organic solutions has been described previously and methods have been developed, the most widely used being UNIversal Functional Activity Coefficient (UNIFAC) (Fredenslund et al., 1975). Though UNIFAC has been noted to have various shortcomings (e.g. Choi and Chan, 2002b; Cruz and Pandis, 2000), numerous attempts to improve its predictive capability for components of atmospheric importance have been made. The difficulty in using these models to predict the hygroscopicity of a mixed organic aerosol (such as ambient organic aerosol) is that detailed input is required (e.g. the abundance of functional groups and the molecular weight of the organic components). Data of this kind are typically not available, nor is this information predicted by large scale atmospheric models.

The hygroscopic properties of SOA obtained by oxidation of anthropogenic and biogenic volatile organic precursors like aromatics or terpenes in smog chambers have also been investigated (Virkkula et al., 1999; Hegg et al., 2001; Shantz et al., 2003; Varutbangkul et al., 2006; Duplissy et al., 2008). Varutbangkul et al. (2006) showed that the SOA hygroscopicity is inversely proportional to the precursor molecular weight and SOA yield. Duplissy et al. (2008) showed that at higher precursor mixing ratios, which results in higher SOA concentrations, the hygroscopicity of the SOA is lower. They proposed that at higher concentration more volatile (and less oxidized) products partition into the particle phase, leading to a decrease of the hygroscopicity of the particle. In field studies, Gysel et al. (2007) and Sjogren et al. (2008) estimated the average hygroscopicity of bulk OA from ambient aerosol hygroscopicity. Cubison et al. (2006) reported, but did not quantify, an increase of the hygroscopicity of ambient aerosol when the OA became more oxidized.

Jimenez et al. (2009) have recently published a relationship between OA chemical composition and hygroscopicity. Here, this relationship as well as the underlying experiments are described in more detail and are complemented by further experiments. We measured the water uptake of aerosol particles at subsaturated RH with a hygroscopicity tandem differential mobility analyzer (HTDMA) (Liu et al., 1978). Par-

allel measurements of the time and mass-weighted chemical composition of non-refractory submicron aerosol (including sulfate, nitrate, ammonium and OA) were performed with an aerosol mass spectrometer (AMS). The hygroscopic growth of OA was either directly determined for pure OA or, for mixtures of organic and inorganic aerosol, deduced from the hygroscopic growth measurements with the Zdanovskii-Stokes-Robinson (ZSR) relation (Stokes and Robinson, 1966) using the chemical composition obtained from the AMS (Gysel et al., 2007).

2 Experimental

2.1 Smog chamber experiments

The experiments were performed in the PSI smog chamber (Paulsen et al., 2005). The PSI smog chamber is a 27-m³ transparent Teflon[©] (FEP) bag suspended in a temperature-controlled housing. The radiation was generated by four xenon arc lamps (4 kW each) selected to simulate the solar light spectrum and natural photochemistry. Before introduction of the precursors the chamber was humidified to 50% relative humidity. SOA was produced from three different precursors: 1,3,5-trimethylbenzene (TMB), α -pinene and isoprene. The precursor concentrations were measured with a proton-transfer-reaction mass spectrometer. Two different sets of experiments were performed; (1) traditional photo-oxidation experiments and (2) ozonolysis of α -pinene followed by oxidative ageing of the reaction mixture with OH radicals. The experimental conditions are summarized in Table 1.

1. In the traditional photo-oxidation experiments an organic precursor and nitrogen oxides (ratio of 2 to 1) were introduced into the chamber and allowed to mix for approximately 30 min before the lights were turned on. The reaction of OH with TMB leads to formation of RO₂ radicals, which rapidly convert NO to NO₂. Ozone is formed from the photolysis of NO₂, and its concentration increases rapidly after NO concentration has dropped. In the case of α -pinene and isoprene the reaction of ozone also plays a role once the ozone concentration becomes high enough. In all these experiments the organic precursor did not react completely away. Thus, besides oxidative ageing of SOA first-generation oxidation products of the precursor were always condensing on the SOA.
2. In the ozonolysis and ageing experiments ozone (200–500 ppb) was first added to the humidified chamber. Then α -pinene at a concentration of 10 or 40 ppb was introduced. After all α -pinene had reacted away by ozone, the reaction mixture was further oxidized by OH radicals in two different ways, called dark OH or OH+light.

Table 1. Experimental conditions for SOA formed in the smog chamber.

Experiment	Precursor [ppb]	NO [ppb]	NO ₂ [ppb]	O ₃ [ppb]	Remarks
1	TMB 150	–	75		traditional photo-oxidation
2	TMB 86	–	43		traditional photo-oxidation
3	α -pinene 180	48	60		traditional photo-oxidation
4	α -pinene 124	31	41		traditional photo-oxidation
5*	α -pinene 10	–	5		traditional photo-oxidation
6	isoprene 1200	300	300		traditional photo-oxidation
7 ¹	α -pinene 10		200–500		ozonolysis
8 ²	α -pinene 40		110–500		ozonolysis
9 ³	α -pinene 10	–	–		OH exposure of mixture exp 7
10 ⁴	α -pinene 40	–	–		OH exposure of mixture exp 8

* Experiment 5 is a combination of two similar experiments where the AMS and HTDMA were not measuring simultaneously.

¹ six similar ozonolysis experiments

² three similar ozonolysis experiments

³ five similar experiments with OH+light or dark OH

⁴ three similar experiments with OH+light or dark OH

For the dark OH, ozone was added to obtain about 500–700 ppb in the chamber. Then tetramethylethylene was continuously flushed into the chamber where it reacted with ozone and produced OH radicals in high yields. For the OH+light, HONO was flushed into the chamber until it reached approximately 10 ppb. Then lights were turned on and photolysis of HONO produced OH radicals. HONO was continuously replenished to provide a semi-constant OH source.

2.2 Field measurements

Field measurements were performed at the high-alpine site Jungfraujoch, Switzerland and in Mexico City.

The Jungfraujoch is a European high-alpine background site located on an exposed mountain ridge in the Bernese Alps, Switzerland, at 3580 m altitude (46.33° N, 7.59° E). Within the World Meteorological Organization (WMO) Global Atmosphere Watch (GAW) program continuous measurements of aerosol parameters have been performed at the JFJ site since 1995 (Collaud Coen et al., 2007). The station is surrounded by glaciers and rocks, and no local vegetation is present. An Aerodyne quadrupole AMS and an HTDMA were deployed at the Jungfraujoch research station. Hygroscopicity measurements have been performed during the Cloud and Aerosol Characterization Experiment (CLACE3) (Sjogren et al., 2008).

The measurements in Mexico City were part of the Megacity Initiative: Local and Global Research Observations (MILAGRO) and took place in and around Mexico City during March 2006. The MILAGRO campaign was designed to study the chemical characterization and transformation of pollutants from the Mexico City urban area

to regional and global scales. A high resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) and an HTDMA were deployed on an aircraft platform onboard the National Science Foundation/National Center for Atmospheric Research (NSF/NCAR) C-130 aircraft. Details on the deployment and results for the campaign are found in DeCarlo et al. (2008, 2010).

2.3 HTDMA

The hygroscopic properties of the various aerosol types were measured with a HTDMA. Briefly, a HTDMA functions as follows: a differential mobility analyzer selects a monodisperse aerosol size cut with mobility diameter, D_0 , under dry conditions. The aerosol then passes through a humidifier with a controlled higher RH, and the size distribution over mobility diameter $D(RH)$ is measured with a second DMA coupled to a condensation particle counter (CPC). The hygroscopic growth factor (GF(RH)) indicates the relative increase in mobility diameter of particles due to water absorption at a certain RH, and is defined as:

$$GF(RH) = \frac{D(RH)}{D_0} \quad (1)$$

In this study three different custom built HTDMAs were used. Specific details of HTDMA1, HTDMA2 and HTDMA3 can be found in Duplissy et al. (2008, 2009), Gasparini et al. (2004) and Sjogren et al. (2008) respectively. The residence times of the aerosol in the humidifier, the chosen RH and the operating temperature were (15 s, 95%, 20 °C), (10 s, 85%, 40 °C) and (20 s, 85%, 25 °C), respectively.

2.4 Aerosol mass spectrometer

An Aerodyne quadrupole AMS was used during type 1 chamber studies and at the Jungfraujoch to provide online, quantitative measurements of the size distributed non-refractory chemical composition of the submicron ambient aerosol at a high temporal resolution. An HR-ToF-AMS (DeCarlo et al., 2006) was deployed on the NCAR C-130 aircraft and type 2 chamber studies. More detailed descriptions of the AMS measurement principles and various calibrations (Canagaratna et al., 2007), its modes of operation (DeCarlo et al., 2006) and data processing and analysis are available elsewhere. The AMS provides the concentrations of inorganic ions, i.e. sulfate, nitrate, ammonium, and fragment ions associated with OA. In addition, the OA mass spectra provide further information on the aerosol. Several mass-to-charge ratios, specifically, m/z 44, m/z 57, and m/z 60 have been proposed as markers for oxygenated species, hydrocarbon-like (mostly urban combustion), and wood burning OA, respectively (Canagaratna et al., 2007; Alfarra et al., 2007). For the correlation analysis with particle hygroscopicity we use in the following the ratio of specific AMS OA mass fragments to total OA mass (f_x , with x being any integer m/z). As an example, for m/z 44 this abundance is defined as

$$f_{44} = \frac{m/z 44}{\text{Total organic mass}} \quad (2)$$

3 Theory and data analysis

3.1 The growth factor GF

As shown in Eq. (1), the GF is associated with a certain RH. The water activity, a_w , of a solution is defined as the equilibrium RH over a flat surface of this solution (i.e., in the absence of the Kelvin effect). The Köhler equation, $\text{RH} = a_w S_k$, describes the equilibrium RH for a solution droplet, where S_k is the Kelvin factor defined as follows:

$$S_k = \exp\left(\frac{4M_w\sigma_{\text{sol}}}{RT\rho_w D_p}\right) \quad (3)$$

where M_w is the molar mass of water, ρ_w is the density of water, σ_{sol} is the surface tension of the solution, R is the ideal gas constant, T is the temperature and D_p is the wet diameter. In this study we used the surface tension of water for σ_{sol} . To compare GF values measured at different a_w , we used the semi-empirical model described in Petters and Kreidenweis (2007), where GF and a_w of a particle are related as follows:

$$\text{GF}(a_w) = \left(1 + \kappa \frac{a_w}{1 - a_w}\right)^{1/3} \quad (4)$$

where the hygroscopicity parameter κ captures all solute properties (such as the number of dissociated ions per molecule and the molal osmotic coefficient). Figure 1 shows

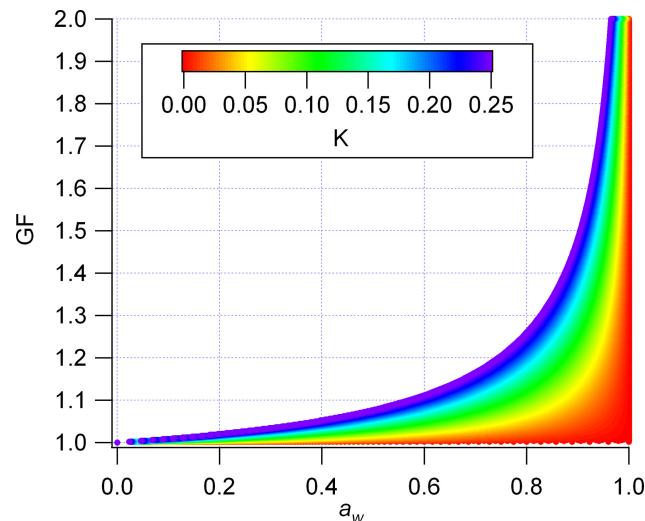


Fig. 1. Influence of the hygroscopicity parameter κ on the relationship between the hygroscopic growth factor (GF) and the water activity (a_w).

the influence of κ on the relationship between the growth factor (GF) and the water activity (a_w). More details about the theoretical background of the functionality of Eq. (4) are given in Kreidenweis et al. (2005) and Petters and Kreidenweis (2007).

3.2 Retrieval of organic GF (GF_{org}) from an internally and externally mixed aerosols

The hygroscopic growth factor of OA was directly measured in the pure SOA smog chamber experiments. However, the hygroscopicity of the organic fraction in ambient aerosol has to be deduced from the GF of the mixture (GF_{mixed}). The GF_{mixed} can be estimated from the growth factors of the individual components of the aerosol and their respective volume fractions, ε , with the ZSR relation (Meyer et al., 2009):

$$\text{GF}_{\text{mixed}} = \left(\sum_i \varepsilon_i \text{GF}_i^3 \right)^{1/3} \quad (5)$$

with the summation performed over all compounds present in the particles. Solute-solute interactions are neglected in this model and volume additivity is also assumed. The concentrations of ammonium (NH_4^+), sulfate (SO_4^{2-}), nitrate (NO_3^-), and OA during these two field campaigns were obtained from the AMS measurements. In both campaigns, the aerosol was mostly neutralized with ammonium (DeCarlo et al., 2008; Cozic et al., 2008). The pairing of the inorganic ions, which is required for the ZSR relation, is unambiguous for aerosols neutralized by a single cation. However, occasionally the measured ammonium concentration was insufficient to fully neutralize the sulfate, thus indicating an acidic aerosol. In such cases the ion-pairing becomes ambiguous

Table 2. GF and density used in the ZSR calculation (Bulk properties; Topping et al., 2005a).

Substance	GF at $a_w = 0.85$	Densities [kg m $^{-3}$]
$(\text{NH}_4)_2\text{SO}_4$	1.56	1769
NH_4HSO_4	1.62	1720
NH_4NO_3	1.59	1780
H_2SO_4	1.88	1830
Organic	–	1270 ^a

^a The density of organic was chosen to represent oxidized organics in aged atmospheric aerosol (Cross et al., 2007).

and therefore an adequate ion-pairing scheme must be applied. We excluded such periods and used the simplified ion-pairing scheme presented by Gysel et al. (2007), which has a direct analytical solution and deviations of corresponding ZSR predictions from full thermodynamic models are minor for the relevant inorganic mixtures. Bulk growth factors of pure inorganic salts were obtained from the Aerosol Diameter Dependent Equilibrium Model (ADDEM; Topping et al., 2005a; Table 2). Densities used to convert mass fractions measured by the AMS into volume fractions required for the ZSR relation (Eq. 5) are also provided in Table 2. The growth factor of OA, the only unknown variable of Eq. (5), can then be calculated from the AMS (providing ε) and HTDMA (providing GF_{mixed}) data.

When several growth factors were measured simultaneously (which occurred in Mexico data) due to external mixture of the aerosol, an average GF was then calculated using the following equation:

$$\text{GF} = \left(\sum_i n_i \text{GF}_i^3 \right)^{1/3} \quad (6)$$

where n_i is the number fraction of particles having the growth factor GF_i . An important source of uncertainty for the retrieved GF_{org} is the organic density used. For ambient organic aerosol the densities reported in literature vary from 1200 to 1700 kg m $^{-3}$ (Hallquist et al., 2009). Another source of uncertainty for the retrieved GF_{org} comes from the OA volume fraction (ε_{org}). As shown in Fig. 2, the GF_{org} uncertainty increases strongly as ε_{org} decreases, since it is increasingly being determined as the difference of two large numbers in Eq. (5). The GF_{org} uncertainty also increases as the OA hygroscopicity decreases, since the size of the $\varepsilon_{\text{org}} \text{GF}_{\text{org}}^3$ term in Eq. (5) is becoming smaller relative to the inorganic and total terms. Therefore, to keep the uncertainty lower, we restricted our data to ε_{org} higher than 80% in the case of non hygroscopic organic ($\text{GF}_{\text{org}}(85\%) = 1$) and 25% for very hygroscopic organic ($\text{GF}_{\text{org}}(85\%) = 1.3$). To keep a low ε_{org} uncertainty, AMS data with a mass loading of OA

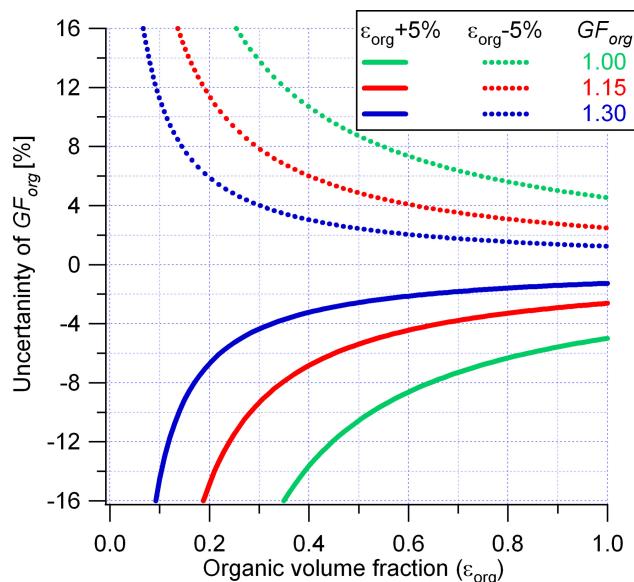


Fig. 2. Uncertainty of the retrieved GF_{org} due to AMS uncertainties. Errors calculated with $\varepsilon_{\text{org}} \pm 5\%$ ($\varepsilon_{\text{org}} - 5\%$ for dashed line and $\varepsilon_{\text{org}} + 5\%$ for the filled lines) and an inorganic $\text{GF} = 1.56$, assuming 3 different GF_{org} : 1, 1.15 and 1.3. With a higher GF_{org} or with a higher ε_{org} , the uncertainty of the calculated GF_{org} decreases. For example, for a non hygroscopic organic ($\text{GF}_{\text{org}} = 1$), the retrieved GF_{org} is 1 ± 0.06 at $\varepsilon_{\text{org}} = 0.8$ and 1 ± 0.14 at $\varepsilon_{\text{org}} = 0.4$, whereas for a hygroscopic organic ($\text{GF}_{\text{org}} = 1.3$) the retrieved GF_{org} is 1.3 ± 0.02 at $\varepsilon_{\text{org}} = 0.8$ and 1.3 ± 0.04 at $\varepsilon_{\text{org}} = 0.4$.

less than 2 $\mu\text{g m}^{-3}$ were excluded. The non-refractory materials (like black carbon and mineral dust) which are not measured by the AMS add additional uncertainty to the field data. These materials are typically non-hygroscopic and therefore the derived GF_{org} can be underestimated. From GF_{org} , the hygroscopicity parameter for OA, κ_{org} was determined using Eq. (4).

3.3 Statistical method for correlation analysis of the smog chamber data

The forward screening method is used to evaluate the correlation of AMS data with κ for the smog chamber data.

A thorough description of this method is given in Wilks (2006). Briefly, in a first step, the linear correlation coefficient of κ with each f_x is calculated and the f_x yielding the highest value is kept for the next step. In the second step, a tri-linear regression analysis with κ and the f_x selected from the first step and with each remaining f_x is performed (i.e. $\kappa = a \times f_{1\text{st step}} + b \times f_x + c$). The f_x yielding the best prediction of κ is then kept with the previous one for the next step. The method is stopped when there is no more significant improvement of R^2 ($\Delta R^2 < 0.05$) when adding another f_x .

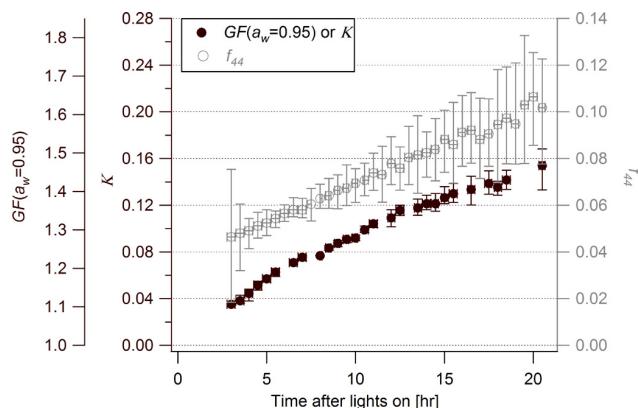


Fig. 3. Temporal evolution of the GF at $a_w = 0.95$ and the corresponding κ_{org} value (closed symbol) and f_{44} (open symbol) for SOA formed from 150 ppb TMB and 75 ppb NO₂ (Experiment 1 from Table 1). Each data point represents a 30-minute average.

4 Results and discussion

4.1 Smog chamber data

In the smog chamber experiments, the SOA GF was measured with HTDMA1. The dry diameter was adjusted according to the number mode of the aerosol during its growth. Figure 3 shows the typical time trends during a photo-oxidation experiment with TMB (experiment 1 in Table 1). Both the hygroscopicity of the SOA (expressed as κ_{org} and GF ($a_w = 0.95$)) and f_{44} increase during the course of the experiment, i.e. with increasing integrated oxidant exposure. A similar trend of κ_{org} and f_{44} was measured for SOA formed from α -pinene (see Fig. 1 in Duplissy et al., (2008)) as well as from isoprene.

For each chamber experiment f_{44} showed the highest linear correlation of any f_x with κ_{org} , (with $R^2 > 0.92$ and significance $\alpha = 0$ excluding experiment 5 where κ_{org} and f_{44} were not measured simultaneously, see Table 1). By comparison, no other f_x was found to correlate with κ_{org} with $R^2 > 0.64$. Figure 4 shows the correlation coefficient of the estimated κ_{org} versus the different f_x for the three different precursors. Both correlation ($R > 0$) and anti-correlation ($R < 0$) are observed. Specific f_x (like for example f_{55}) can either correlate or anti-correlate with κ_{org} depending on the precursor. In addition, the R between κ and each f_x for all smog chamber data (all precursors together) with significance lower than 0.05 is shown in Fig. 5. With all experiments combined together, a correlation analysis of the calculated κ_{org} and f_x has been performed using the screening regression method and using f_{30} , f_{44} and f_{46} (as suggested by Fig. 5) as three different initial first steps. In doing so, these three analyses suggest that f_{30} , f_{44} , f_{46} and f_{55} are the four parameters which should be used to make a multiple linear regression with κ_{org} (see Fig. 6). It should be noted that m/z 30 and 46 for laboratory data are assumed to be due

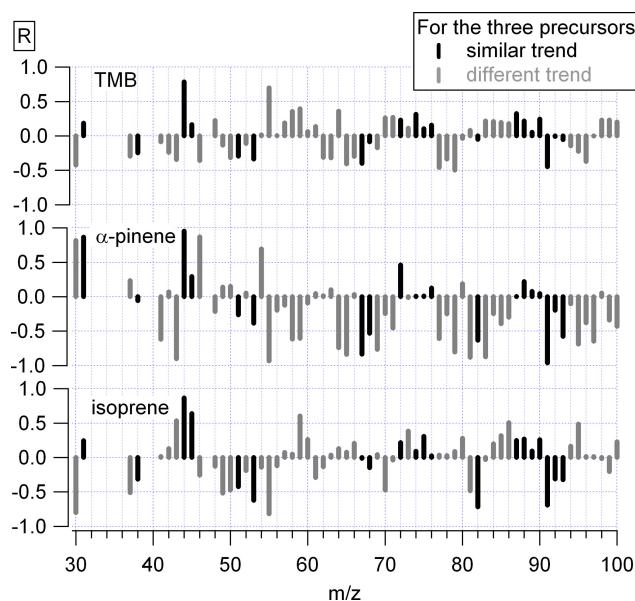


Fig. 4. Correlation analysis of the estimated κ_{org} with f_x (for m/z 30 to m/z 100) for the photooxidation experiments with different precursors. Black bars indicate that the m/z had a similar sign for R for all precursors studied here. In contrast grey bars indicate that for one precursor R has a different sign compared to the other precursor for a same fragment.

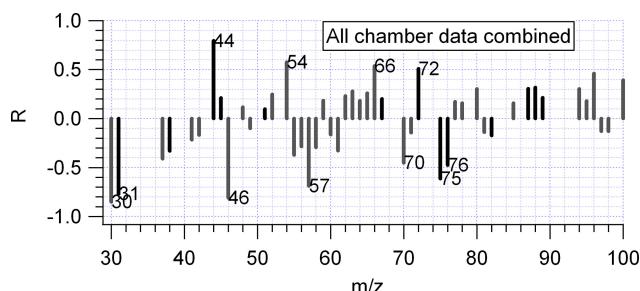


Fig. 5. Correlation coefficient between κ_{org} and each f_x for the complete smog chamber data set. Black bars indicate that the m/z had a similar sign for R for all precursors studied here. In contrast grey bars indicate that for one precursor R has a different sign compared to the other precursor for the same fragment. Data with a significance $\alpha > 0.05$ are not shown.

to OA, even if the signals arise also from NO⁺ and NO₂⁺. This assumption is not valid for ambient data where m/z 30 and 46 are typically dominated by inorganic ammonium nitrate (Zhang et al., 2007b), and as such these signals are not dominated by variations in the composition of the OA.

From a chemical point of view f_{43} ($f_{43} > 15\%$) could also have a correlation with GF from SOA generated in smog chambers and could also be considered as an additional parameter to improve the correlation. However, as shown in Fig. 4, f_{43} and GF can be either correlated or anti-correlated depending of the precursor. For the full data set, adding

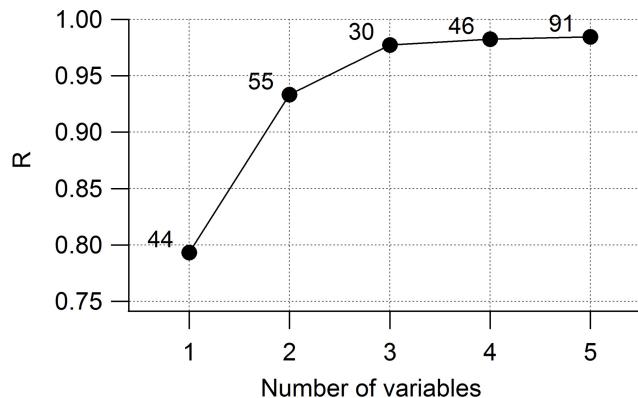


Fig. 6. Result from the screening multiple linear regression of all smog chamber data sets, with f_{44} as the first input variable. Choosing f_{30} or f_{46} or f_{55} as first input variable for the screening regression results in the same four f_x being selected by the procedure (not shown).

f_{43} to f_{44} improves the correlation with κ_{org} (R^2 increases from 0.8 to 0.9). However, ambient aerosol contains not only SOA but also a substantial contribution from primary OA (POA), and both m/z 43 and m/z 55 are prominent in reduced POA mass spectra (typically from the ions C_3H_7^+ and C_4H_7^+ , respectively) (Lanz et al., 2007; Canagaratna et al., 2004; Zhang et al., 2005), besides being components of the SOA mass spectra (typically the ions $\text{C}_2\text{H}_3\text{O}^+$ and $\text{C}_3\text{H}_3\text{O}^+$, respectively) (Alfarra et al., 2006). Due to the different sources of m/z 43 and m/z 55, and the differing hygroscopic properties of OA associated with these sources, ambient datasets are less suitable than chamber data for fitting κ_{org} with these m/z . m/z 44 is mainly associated with the CO_2^+ ion fragment, and data from the HR-ToF-AMS indicate that there is only a minor contribution from $\text{C}_2\text{H}_4\text{O}^+$ in SOA as well as in ambient air (Mohr et al., 2009). Therefore, due to these interferences on ambient OA measurements, only f_{44} can be used here in both lab and field data as a proxy for κ_{org} .

The correlation between κ_{org} and f_{44} is shown in Fig. 7 and is also found in field experiments (see below). It is likely explained by the fact that m/z 44 is a major fragment of highly oxidized organic species such as organic di-acids, poly-acids, oxo-acids, hydroxy-acids, and acyl peroxides (Aiken et al., 2007; Aiken et al., 2008), and is also highly correlated with the bulk atomic O/C ratio of the OA (Aiken et al., 2007; Aiken et al., 2008), which is given as a secondary x-axis in Fig. 7. These functional groups are expected to have a rather high hygroscopicity. Plotting f_{44} versus κ_{org} the data lies on the same line for α -pinene and isoprene SOA, while data from TMB SOA are shifted to higher κ_{org} for the same f_{44} (Fig. 7). This result may be due to the fact that in SOA from TMB photooxidation, the acids present in the aerosol are mainly mono-acids (> 90%), while the acids found in the SOA from α -pinene or isoprene photooxidation are approximately half mono- and half di-acids (Gäggeler, 2008).

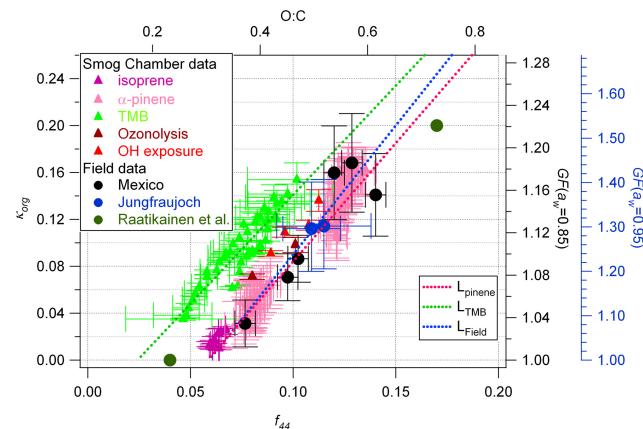


Fig. 7. Relationship between f_{44} and κ_{org} for field data (Jungfraujoch and Mexico City), smog chamber data (SOA formed from traditional photooxidation of isoprene, α -pinene and 1,3,5-trimethylbenzene, from the ozonolysis of α -pinene and subsequent OH exposure. Data of Raatikainen et al. (2010) are also shown. Linear regression fits are shown for TMB (L_{TMB}), isoprene/ α -pinene (L_{pinene}) and field data (L_{field}).

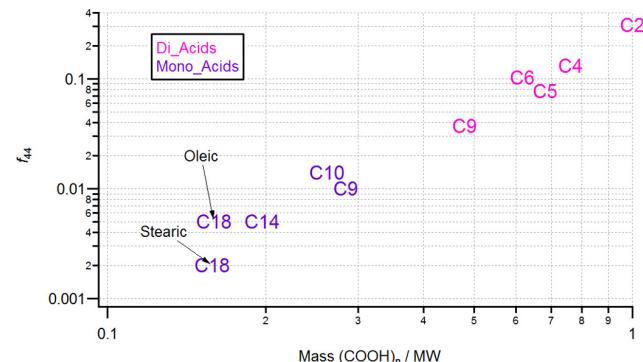


Fig. 8. The fractional contribution of m/z 44 (f_{44}) for mono- and di-carboxylic acids versus the ratio of the mass of carboxylic acid groups to the molecular mass of the individual acids. This figure normalizes the mass of the acid groups to the molecular weight and compares it to f_{44} in the mass spectrum for each acid. It shows a tendency for di-acids to produce more m/z 44 per COOH group than mono-acids.

Alfarra (2004) using pure standards showed that the signal response of m/z 44 per organic acid group is lower for mono-acids than for diacids (Fig. 8). This is consistent with the offset observed here in Fig. 7 between TMB and biogenic precursors.

4.2 Ambient data

The aerosol at the Jungfraujoch shows an annual cycle with highest mass concentrations in July to August and minimum concentrations in January to February. The OA fraction is higher during the summer season (e.g. Cozic et al., 2008).

Periods with OA concentrations higher than $2 \mu\text{g m}^{-3}$ and a volume fraction of OA higher than 80% were only found during the summer CLACE3 campaign, and therefore only data from this period are shown. The restriction of the data set to these mass loadings and volume fractions is based on the uncertainties to calculate GF_{org} as described above. 3-h averages were calculated as a trade-off between counting statistics and time resolution. HTDMA2 measured the hygroscopic properties of particles with $D_0 = 100 \text{ nm}$. The AMS composition data were integrated over a narrow size range from 88–196 nm vacuum aerodynamic diameter, which corresponds to a volume-equivalent diameter of 100 nm (56–125 nm) as described in Sjogren et al. (2008) and using the framework detailed in DeCarlo et al. (2004).

A large fraction of the OA in Mexico City is oxidized, with an oxygen-to-carbon atomic ratio (O/C) that increases with distance from the city approaching a ratio of 0.9 away from the basin indicating photochemical ageing of OA (DeCarlo et al., 2008; Aiken et al., 2008). The NCAR-NSF C-130 performed 12 research flights (RFs) during the MILAGRO campaign. Measurements in the aircraft constrain the black carbon (BC) and dust fractions to be small, of the order of a few percent of the total submicron mass (DeCarlo et al., 2008). Data reported here are from RFs 3 and 12 (10 and 29 March 2006, respectively) and both had city and regional components in their flight tracks (DeCarlo et al., 2008, 2010). The OA contained important influences from urban POA, anthropogenic SOA, and biomass burning POA and SOA (DeCarlo et al., 2010). The aerosol mass distribution was dominated by larger particles ($D_{\text{va}} > 200 \text{ nm}$). Therefore, all AMS data were integrated (no size distinction) and GFs only from larger particles ($D_{\text{m}} = 200 \text{ nm}$ and 300 nm) measured by HTDMA3 were analyzed assuming the same bulk chemical composition. A limitation of aircraft measurements is that air masses can change rapidly, often within an HTDMA measurement cycle. Therefore care was taken to include only data where the aerosol mass and composition were stable during an HTDMA3 scan. This was determined using the AMS data (12 second resolution), and data with a variation of either the f_{44} or the organic to inorganic ratio higher than 10% during the hygroscopicity measurement were excluded.

For each period with a valid calculation of GF_{org} and corresponding κ_{org} , f_{44} was extracted from the AMS measurements. Ambient data are also presented in Figure 7. It is obvious from this figure that κ_{org} and f_{44} show a strong and consistent dependence and that κ_{org} may be approximated from f_{44} according to the equation

$$\kappa_{\text{org}} = a \times f_{44} + b \quad (7)$$

The parameters a, b are given in Table 3 for this relationship for TMB, α -pinene and isoprene as well as the ambient data separately. Smog chamber data from both biogenic precursors and field data at both locations show a consistent relationship between f_{44} and κ_{org} within the uncertainties. The

Table 3. Correlation parameters of Eq. (7) for the different aerosols.

Variable	a	b
TMB	1.86 ± 0.02	-0.045 ± 0.01
α -pinene, isoprene	2.02 ± 0.04	-0.118 ± 0.005
Jungfraujoch/Mexico	2.2 ± 0.4	-0.13 ± 0.05

slope for TMB is similar to the other results but there is an offset as explained above. Recently Raatikainen et al. (2010) reported growth factors of 1.0 and 1.29 at 88% RH for 50-nm particles (corresponding to κ_{org} of 0 and 0.2, respectively) at f_{44} of 0.04 and 0.17 respectively. Both values are in good agreement with our data and also support our observation that OA with $f_{44} < 0.06$ seems to be effectively non-hygroscopic.

In this study we investigate the relationship between the oxidation level of the SOA and its hygroscopic growth factor at subsaturated RH. Chang et al. (2010) reported a similar relationship for κ_{org} ($\kappa_{\text{org}} = (0.29 \pm 0.05) \times (\text{O/C})$) derived from ambient cloud condensation nuclei (CCN) measurements at supersaturated RH. This slope differs from our results, which can likely be attributed to decreased importance of non-ideal interactions in dilute solutions at the point of CCN activation (Petters et al., 2009). Several laboratory studies reported differences between HTDMA- and CCNC-derived κ_{org} of SOA (e.g. Prenni et al., 2007; Wex et al., 2009; Juranyi et al., 2009, Good et al., 2010). Massoli et al. (2010) showed for different proxies of anthropogenic and biogenic SOA that the relationship between O/C ratio and CCN-derived kappa value is non-linear.

5 Conclusions

A simplifying approximation, such as the one proposed here, could introduce an important improvement in the parameterization of hygroscopicity in atmospheric models, since f_{44} is significantly correlated with the photochemical age of the air mass (DeCarlo et al., 2008; Aiken et al., 2008) and the database for such AMS measurements is constantly being expanded around the world (Jimenez et al., 2009). We recommend testing the general applicability of Eq. (7) with more ambient and chamber measurements, including a variety of different POA sources (e.g. diesel exhaust and biomass burning OA, fresh and aged), and in both the sub- and supersaturated regimes. The ability of the HR-ToF-AMS to distinguish individual ions at the same nominal m/z could potentially eliminate some of the interferences mentioned above for other nominal m/z and yield other ions that can be used to improve the κ_{org} calculation. If generally applicable, this relationship would substantially improve our ability to describe the evolution of hygroscopicity at subsaturated RH in a variety of atmospheric models.

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